

## BY-PASS SEDIMENT SYSTEM TO REPLACE LONGSHORE CURRENT IN HARBOUR AREAS

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### EXTENDED ABSTRACT

Il bilancio sedimentario a scala del sistema bacino idrografico-area di costa è molto importante nel sistema di gestione sia delle aree fluviali che di quelle costiere. Nell'ambito del progetto Tech4You (T4Y S2G2PP2), è stata analizzata la problematica dell'insabbiamento dell'imboccatura del porto di Cetraro (Calabria, Italia meridionale). In merito al clima meteomarinario, l'unità fisiografica che include il porto di Cetraro è direttamente esposta all'azione di venti provenienti da sud-ovest, ovest e nord-ovest attraverso il Mar Tirreno. Il vento dominante soffia da ovest-sud-ovest con un fetch geografico maggiore di 1650 km e velocità fino a 7 Beaufort ( $\geq 28$  nodi). Queste velocità sono frequenti in inverno e producono le onde più alte fino a 5m. I venti prevalenti sono quelli che provengono da nord-ovest, con un fetch massimo di 650 km e velocità minori di 6 Beaufort ( $< 22$  nodi). Questi venti soffiano prevalentemente durante la primavera e l'estate. Il movimento longshore dei sedimenti, che avviene da nord verso sud, è interrotto dal molo di sopraflutto del porto che causa un consistente processo di accrezione della spiaggia sul lato Nord del porto e insabbiamento alla sua imboccatura ( $\sim 10,000$  m<sup>3</sup>/anno di materiale depositato); al contrario, nel tratto di costa sottoflutto, sono stati osservati intensi processi di erosione verosimilmente dovuti alla forte riduzione di apporto sedimentario lungo costa.

Nell'ambito del bilancio sedimentario dell'unità fisiografica, è stata analizzata la produzione di sedimento all'interno del bacino idrografico del Fiume Aron. Quest'ultimo rappresenta la principale sorgente di alimentazione sedimentaria all'interno della stessa unità fisiografica. La potenziale produzione media annua di sedimento del bacino idrografico è stata stimata attraverso il Modello dell'Erosione Potenziale (EPM) di Gavrilovic, utilizzando il plug-in PQGIS YES che permette di applicarle l'EPM in modo semi-automatico. L'EPM è un modello semiquantitativo che si basa su dati geologici, geomorfologici, climatici e di uso del suolo. Partendo da questi dati, il modello permette di stimare la potenziale produzione di sedimento del bacino idrografico, il trasporto e la sedimentazione.

L'uso del suolo all'interno del bacino del Fiume Aron è stato mappato mediante l'analisi di immagini satellitari multitemporali e multispettrali utilizzando tecniche machine learning, come l'algoritmo "Random forest". I dati ottenuti sono stati successivamente validati da specifiche verifiche di campo. I principali risultati dell'analisi multitemporale hanno mostrato un incremento nel tempo delle aree boschive ai danni delle aree seminative e dedite al pascolo, correlato al progressivo abbandono delle aree rurali. Lungo il settore costiero è stato invece osservato un incremento delle aree urbanizzate. Il coefficiente relativo alla litologia, previsto nel modello EPM, è stato attribuito ai diversi litotipi affioranti nel bacino dopo averne valutato la resistenza attraverso specifici rilievi geomeccanici. Per il bacino del Fiume Aron, la stima del volume medio di sedimenti prodotti annualmente ottenuta con il modello EPM risultata pari a  $\sim 30,000$  m<sup>3</sup>/anno.

La quantificazione dei volumi di sedimento coinvolti nel bilancio sedimentario costiero e la conoscenza delle correnti nel settore sotto costa ha permesso di ideare una soluzione progettuale in grado di mitigare allo stesso tempo l'insabbiamento dell'imboccatura del porto e i processi di erosione nell'area sottoflutto. Nel dettaglio, è stato progettato un sistema di *bypass* consistente in una pompa sommersa, da installare nella porzione finale del molo nel settore sopraflutto, avente lo scopo di dragare autonomamente il materiale sedimentario accumulato e riversarlo nel tratto di costa a sud del porto. Il sistema di *bypass* permette quindi di movimentare autonomamente il sedimento dalla trappola generata dalla struttura portuale nel settore costiero sottoflutto. In pratica, il sistema permette di riattivare artificialmente la naturale corrente di longshore interrotta dalla costruzione del porto. Contemporaneamente è stato sviluppato anche un sistema di monitoraggio che permette di valutare l'impatto del funzionamento del *bypass* sulle diverse matrici ambientali.

## ABSTRACT

This paper is part of the results achieved in Tech4You project T4Y S2G2PP2 funded by PNRR (Piano Nazionale di Ricerca e Resilienza). The Tech4You project focused on the silting process affecting the Cetraro harbour mouth (Calabria, southern Italy). The study area is located along the northern Calabrian Tyrrhenian coast and is part of a physiographic unit with a west-southwest fetch greater than 1650 km. The wave climate is characterised by a significant wave height of 0.65m (average period of 5.15s) and 1.01 m (average period of 5.7s) offshore and inshore, respectively. The longshore flow of sediments (about 80,000 m<sup>3</sup>/yr), moving from north to south, is interrupted by the harbour pier producing silting (~10,000 m<sup>3</sup>/yr sediment deposited at the harbour mouth), and heavy coastal erosion in the down-drift sector. The sediment production in the Aron River catchment (representing the main sediment feed in the physiographic unit) by Gravrilovic model was also studied, with detailed analysis of soil use through multispectral analysis, and estimate an average volume of ~30,000 m<sup>3</sup>/yr. A bypass system was designed, consisting of a submerged pump installed at the end of the pier on the up-drift side of the harbour, to replace the original longshore current, contrasting simultaneously both mouth silting and erosion processes in the down-drift coastline.

**KEYWORDS:** sediment management, dredging, sustainable harbours.

## INTRODUCTION

Coastal erosion currently affects much of the worldwide Nation, in particular areas characterised by the presence of widespread anthropic structures (e.g. ANFUSO & DEL POZO, 2005; STANCHEVA *et alii*, 2011; IETTO *et alii*, 2018).

Anthropogenic pressure is a serious problem affecting many coastal areas (e.g., ALVAREZ-CUESTA *et alii*, 2021; IETTO *et alii*, 2014).

The coastal anthropization area has increasingly accelerated in the second half of the twentieth century (e.g., IETTO *et alii*, 2014; CANTASANO *et alii*, 2017). In fact, after the end of the Second World War, a considerable migration from inland to coastal territories was observed in most coastal areas in the Mediterranean region (ROMANO *et alii*, 2017). It is significant for example, the number of coastal cities, that is quintupled in the last 70 years, and the percentage of the population living close to the coastal areas which is about 20% of the world population (e.g. STRONKHORST *et alii*, 2018)

The anthropogenic pressure produced a disequilibrium in both coastal and river environments, that represents the trigger of the shoreline erosion processes (e.g. BOMBINO *et alii*, 2022; FOTI *et alii*, 2022). On coastal areas it mainly includes the construction of new settlements (often substituting beaches and dune systems), expanding existing coastal towns, and the construction of port and coastal structures (e.g., CANTASANO *et alii*, 2023). Furthermore, in the fluvial environments,

anthropogenic structures such as dams and weirs represent traps for the bedload transport, producing consequently a deficit in the sedimentary balance of the littoral areas.

FOTI *et alii* (2022) analyzed urban expansion along the Calabrian coasts during the last 70 years, showing an increase of urban centers from 32 to 83, with an increase of urbanized area from 15 km<sup>2</sup> to 250 km<sup>2</sup>. The same authors, in according to previous researchers (D'ALESSANDRO *et alii*, 1998, 2002; IETTO, 2001; GUIDUCCI & PAOLELLA, 2004), highlighted that the increase of urbanisation triggered widespread coastal erosion phenomena. In detail, the main erosive effects are recorded along the northern Tyrrhenian coast (Figure 1a,b), which represents the pilot area of this study. In particular, this research studied the coastal area surrounding Cetraro village in Cosenza province (southern Italy, Figure 1c), where the morphological evolution was investigated through the scientific literature. In detail, IETTO (2001) reconstructed and interpreted the shoreline evolution of the northern Tyrrhenian Calabrian coast between 1953 and 1999, showing a shoreline retreat rate of 0,9±1 m/yr in the period 1953-1983, with a deceleration in the subsequent period 1983-1999. Referring to the Cetraro Marina area, the same Author showed a shoreline retreat of 120 m in the period 1953-1983, followed by an advance of 12 m in the period 1983-1999, he highlights a shoreline advance of 40 m, during the period 1954-1983, in correspondence with the Cetraro harbour up-drift side. Furthermore, IETTO *et alii* (2018) argued on a new coastal erosion risk assessment indicator applied to the northern Tyrrhenian sector. They showed that 35% of the coastal sector could be classified as very high risk, 30% as high risk, 28% as medium risk and only 7% as low risk.

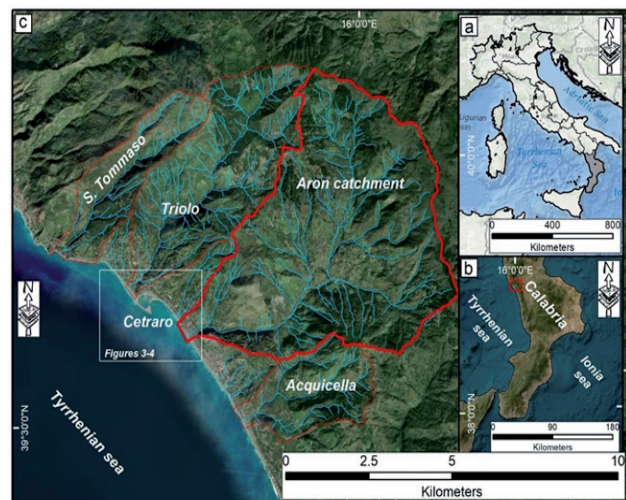


Fig. 1 - Location of the study area. a) location of Calabria (greyish colour) region in southern Italy; b) Location of the studied area in the northwestern part of the Catena Costiera Calabria (red rectangle); c) Bold Red polygon represents the Aron River catchment and in light blue lines its tributaries, also represented the adjacent S. Tommaso Triolo and Acquicella catchments

In this context, the present research focuses on the silting process affecting the Cetraro harbour mouth. In detail, the aim of the study, part of the Tech4You project (T4Y S2G2PP2), consists of the design of a bypass system that allows the movement of sediments from the artificial obstruction created by the harbour structure to the down-drift coastline. Usually, the consolidated method to remove the deposited material in the harbour mouth consists of the dredging process, allowing the correct navigation. However, this process has some drawbacks: (i) relatively high and low predictable costs; (ii) significant impact on marine flora and fauna; (iii) mobilization of pollutants already present on the seabed. This research proposed a technological solution consisting of a permanently submerged pump close to the harbour mouth, allowing more efficient management of sediments which are moved to the down-drift coastline.

**GEOLOGICAL SETTING**

The study area is located along the northern Tyrrhenian coast of the Calabria region, between the mouth of Noce and Savuto rivers, which flow northward and southward, respectively.

The studied littoral sector extends mainly at the foot of the Pollino Massif (north-northward) and Catena Costiera (southward) (Figure 2).

The surrounding area’s geological setting is mainly represented by different tectonostratigraphic units belonging to different paleographic domains (Figure 2). Seven tectonostratigraphic units outcropping in the northern Calabria region, from the bottom to the top: i) Lungro-Verbicaro unit consituted by Meso-Cenozoic metamorphic carbonatic and pelitic sedimentary successions (IANNACE *et alii* 2007); ii)

Pollino-Ciagola units that overlay the Lungro-Verbicaro unit, and it is mainly rappresented by a typically Apenninic carbonate platform succession of the late Triassic to Cretaceous dolostone and limestone covered in discordance by Paleogenic and Neogenic sedimentary succession locally intruded by Miocenic pillow lavas (AMODIO-MORELLI *et alii*, 1976; IANNACE *et alii*, 2007); iii) Ligurian unit, that consist of ophiolitic sequences and deep-sea deposits derived from the western Thethys oceanic basin (AMODIO-MORELLI *et alii*, 1976; FILICE *et alii*, 2015); iv) Castagna unit that consists of a ductiles shear zone of rocks between the contact of Liguride unit and the base of the upper overlaid Sila unit (AMODIO-MORELLI *et alii*, 1976; FILICE *et alii*, 2015); v) Sila unit consist in the continental units, representative of both lower and upper levels of the continental crust and small volumes of subcontinental upper mantle rocks of the Variscan orogeny (LIBERI *et alii*, 2011; PILUSO & MORTEN, 2004; PILUSO *et alii*, 2000); vi) Miocenic succession consist in a clays, arenites and conglomerates of Serravallian to evaporites deposits of Messinian ages (MATTEI *et alii*, 2002); vii) Plio-Quaternary deposits mainly represented by different cronhostratigraphic sequences composed by a clays, sands and conglomerates of fan deltas deposits, marine and fluvial terraces deposits, until eluvio-colluvial and by stream river and beach deposit of Holocene (BROZZETTI *et alii*, 2017).

The actual physiography of the northern Calabria is guided by the activities of the Quaternary and active faults that drive the uplift of the structural high as Pollino massif, like the horst of the Catena Costiera and the Sila Massif, and the subsidence of the Crati and Paola basins (BROZZETTI *et alii*, 2017; CIRILLO *et alii*, 2022). The combining of the uplift, the weathering, the erosion and the landslides and rockfall play a fundamental role in terms of sediment production for the subsequent transportation from the riverbed to the beach in the Calabria shoreline areas (CIANFLONE *et alii*, 2021; TANGARI *et alii*, 2021; IETTO *et alii*, 2022; CIRILLO *et alii*, 2024).

The study area is included in the Paola physiographic unit (LISTI *et alii*, 2010). This is delimited by Capo Bonifati and Capo Suvero promontories northward and southward respectively. This physiographic unit has a length of about 150 km and a depth of closure between -7.6 and -9.9 m a.s.l. The study area can be included in the northern subsector of the Paola physiographic unit, delimited to the north and south by Capo Bonifati and Intavolata promontories respectively. This sector is characterised by longshore dowdrift from north to south and manly coarse sediment supply related to S.Tommaso, Triolo, Aron and Acquicella streams (Fig. 1c) (CNR, 1997).

**METHODS**

The design of the bypass technological solution needed a multidisciplinary approach involving different technical skills. In fact, the different topics analysed in this research

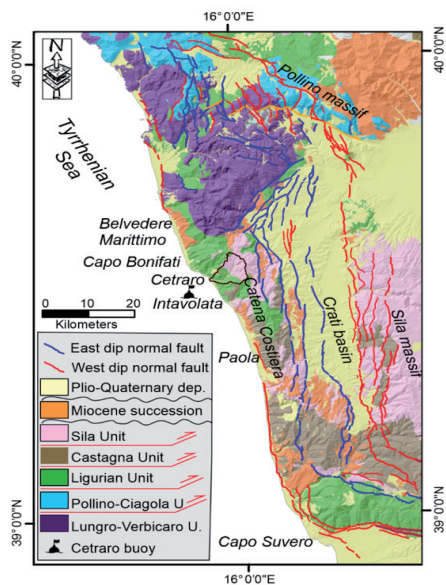


Fig. 2 - Geological map of the northern Calabria modified after Geological Map 25,000 CasMez Calabria (tectono-stratigraphy) and from LAVECCHIA *et alii*, 2024 (Quaternary Faults)

included: (i) wave and current regime; (ii) morphological and sedimentological analysis of the beach sediments in backshore and offshore areas; (iii) sediment balance (including longshore sediment transport and fluvial sediment discharge); (iv) environmental impact of the sediment movement referring to the Italian legislation; (v) submerged pump technology.

The analysis of wave climate was based on wave dataset (from 1999 to 2008) acquired by the Cetraro Buoy (lat. 39.453, long. 15.918), belonging to the Italian National Sea Wave Measurement Network (RON - BENCIVENGA *et alii*, 2012). The dataset collected by the buoy includes: significant wave height, period, and direction at intervals of 30 min.

The analysis of past storm surges identified significant wave heights. The meteorological analysis focused on surface wind, using data from the Belvedere Marittimo and Paola stations.

The analysis of the morphological evolution of submerged areas was based on multi-temporal maps (1954), aerial photos (1985, 1998, 2008) and satellite image (2023). For the submerged area, a multibeam survey has been realized by means of the survey boat of the Marine Laboratory DiBEST-SILA (Unical) equipped with a Multibeam Echosounder Norbit WBMS Basic (400kHz).

The study of littoral transport was based on the calculation realized in the Coastal Master plan of the Calabria Region. This calculation was performed using the model UNIBEST CL+ developed by Delt University. The model, which is not discussed in depth, takes into account different elements: (i) altitude profile of the beach; (ii) wave climate and marine currents; (iii) parameters of wave transformation and breaking; (iv) grain size sediment parameters (D50 and D90). In the same masterplan, the sediment balance was calculated considering: (i) the fluvial sediment discharge; (ii) the littoral transport; (iii) beach nourishment and sand extraction; (iv) offshore sand loss. Starting from this calculation of sediment balance, only the volume of fluvial sediment discharge was modified (see later).

The estimation of the sediment balance in the littoral cell took into account the volume of sediments deposited in the nearshore area (estimated by multibeam survey) and fluvial sediment discharge. For the estimation of the sediment yield and erosion intensity at the basin scale, the erosion potential method (EPM) was used. The EPM (GAVRILOVIC, 1988) estimates the average annual specific production of sediments ( $W$ ) in  $m^3/yr$  through the following equation:

$$W = Th\pi S\sqrt{Z^3}$$

Where  $T$  is the temperature coefficient calculated from the mean annual air temperature ( $^{\circ}C$ );  $h$  is the mean annual precipitation (mm/years);  $S$  is the watershed area ( $km^2$ );  $Z$  is the erosion coefficient:

$$Z = XY(\gamma + \sqrt{Im})$$

Where  $X$  is the land use coefficient;  $Y$  is the coefficient of the rock and soil resistance (a function of geology and soil type);  $\gamma$

is the coefficient of type and extent of erosion;  $Im$  is the average slope steepness of the watershed (%).

The EPM analysis provides an estimated sediment production for the entire catchment area (or at the closure points of individual river basins), but it does not indicate the sectors with different sediment production. To address this, the YES plug-in (DOMINICI *et alii*, 2020) was employed based on GIS techniques and the subdivision of the catchment into a grid matrix. The “squared cell” method calculates sediment production by applying an algorithm to each individual cell.

Different rasters were therefore created containing different information representing the main parameters entering the model and the estimation of sediment yield performed.

In this study, the coefficients of soil resistance to erosion ( $Y$ ) were assigned using the methodology proposed by ZEMPLJIC (1971). The values were obtained from the Calabria geological map at the scale 1: 25,000.

The coefficient of type and extent of erosion ( $\gamma$ ) was obtained using the methodology proposed by ZEMPLJIC (1971) on the data from PAI (the Hydrogeological Plan of southern Italy, <https://gn.mase.gov.it/portale/wfs>).

The average slope of the watershed ( $Im$ ) was calculated using a DEM of the basin area, with cell size of  $20 \times 20$  m ([http://wms.pcn.minambiente.it/wcs/dtm\\_20m](http://wms.pcn.minambiente.it/wcs/dtm_20m)), reclassifying the values into five categories between 0 and 1 (GAVRILOVIC, 1988).  $T$  and  $h$  were obtained from the ISPRA database. ([https://www.isprambiente.gov.it/pre\\_meteo/idro/BIGBANG\\_ISPRA.html](https://www.isprambiente.gov.it/pre_meteo/idro/BIGBANG_ISPRA.html)).

Subsequently, the average annual specific production of sediment ( $W$ ) was calibrated with an innovative approach based on the calibration of the land use parameter through multispectral satellite data processing techniques.

To obtain an accurate estimate of sediment production, the land use coefficient ( $X$ ) was analyzed using highly updated multispectral data and dividing the values into seven categories (ZEMPLJIC, 1971). The image processing was carried out using the particular calculation algorithm Random Forest (RF) (HO, 1995; BREIMAN, 2001). This classifier is an aggregation of a set of decision trees and is based on two fundamental principles: bagging and random subspace. During the construction of the RF classifier, it is therefore necessary to set various parameters such as the number of decision trees and the number of elements on which to train the algorithm.

The procedure involves the use of a supervised classification in which different training areas or Regions Of Interest (ROIs) are identified with which the algorithm is oriented. More than 100 different ROIs were identified, divided by single spectral signature and processed using the SCP plug-in (CONGEDO, 2021) present in the QGIS platform.

The scheme with which the entire procedure was developed provides for the extraction of individual spectral signatures, for each land use category, and their subsequent use in the final

classification. The coefficients calculated as previously described were used to carry out the calculation through the EPM model.

The potential environmental impact of the proposed technical solution was investigated by means of: (i) sediments sampling in both emerged and submerged sites; (ii) 2 multiparametric seawater column logs including measurements of: temperature, pH, turbidity, dissolved oxygen concentration, salinity, density, transparency, fluorescence, and suspended solids. The logs were acquired during May and October 2023 using the multiparametric probe Ocean Seven 316Plus (Idronaut). The turbidity measurements were collected following the protocols: a) method 180.1 of United States Environmental Protection Agency (EPA) (1993) and (b) American Public Health Association - American Water Works Association - Water Environment Federation - Standard Methods for the Examination of Water and Wastewater (1999).

The four sediment samples were subjected to sedimentological identification and characterization and ecotoxicological (referring to APAT-ICRAM, 2007) analysis.

**RESULTS AND DISCUSSION**

To analyse the morphological evolution of the coastal area, multi-temporal coastlines from geographic maps, aerial photos and satellite images were compared (Figure 3). The first comparison is between 1958 (from the Casméz geographic map) and 1985 (from FS-Casméz aerial photo), showing in the up-drift sector with respect to the harbour structure, a coastline increasing with a local maximum of 35 m. Instead, the coastline retreated up to 100 m in the down-drift sector. Between 1985 and 1998 (aerial photo ITA2000 fly), the up-drift sector was characterised by tracts with metric coastline progress and retreat, except the coast stretch close to the harbour breakwater where the coastline progressed up to 50 m. During the same period, also the down-drift sector was characterised by a coastline progress up to 40 m. Between 1998 and 2008 (from aerial photo of Calabria Region), in the up-drift sector 2, coast portions characterised by retreat (up to 25 m) in the northward area and progress (up to 25 m) in the southward were observed. The down-drift sector is instead characterised by coastline retreat up to 17 m. During the period 2008-2023 (from Sentinel-2 satellite image) a general coastline retreat (up to 30 m) in the northern and up-drift sector was observed. Local coastline retreat was also observed in the down-drift sector. Close to the harbour mouth was estimated a deposition of about 10,000 m<sup>3</sup>/yr of sediments.

In the studied area, the evolution of the coastline was hardly influenced by anthropogenic activities. In fact, in the period 1959-1985, the coastline progress and retreat, respectively, in the up-drift and down-drift sectors (with respect to the harbour) can be related to the Cetraro harbour construction that happened during the 1960s. From 1985 to 1998, the relevant coastline progress close to the northern breakwater can be related to the west pier prolongation (realised to contrast the mouth siltation).

Instead, the local coastline progress in the downdrift sector can be ascribed to the realisation of coastal defence measures (beach nourishment and longitudinal emerged breakwaters). Between 1998 and 2008, the up-drift sector continued the coastline progress, while in the down-drift one the coastline retreat persisted despite the realization of new longitudinal emerged breakwaters. Finally, the general coastline retreat can be ascribed to a deficit in the sediment balance due to less solid fluvial discharge related to the variation of land use in the river catchment (see later) and the realisation of hydraulic works along the drainage network.

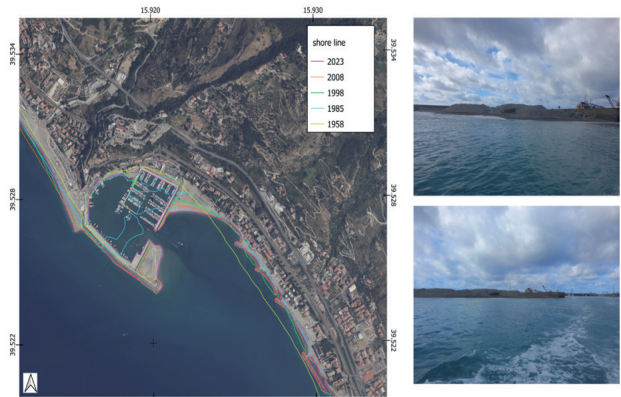


Fig. 3 - Multi-temporal coastlines and photo of deposits at the harbour mouth

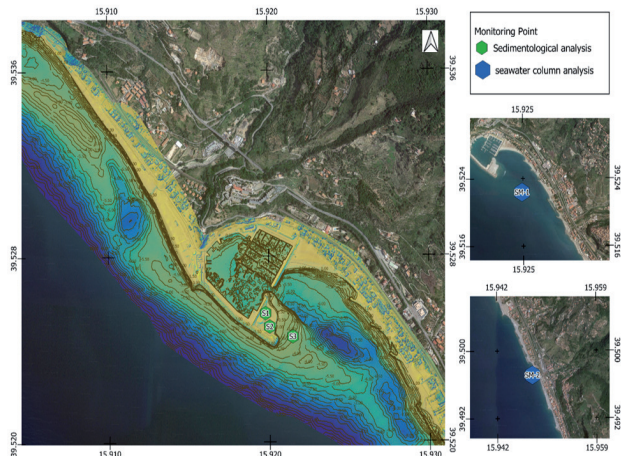


Fig. 4 - Multibeam survey and location of sediments samples (S1, S2, S3) and seawater column multiparametric logs (SM1, SM2)

The multibeam survey allowed to obtain a 1×1 m grid (Figure 4). In the northern sector, up-drift with respect to the harbour, a northwest-southeast bar with a height up to 1.5 m (from -4 to -2.5 m a.s.l.) can be observed. This bar is separated from the coastline by a through deep up to -5.5 m a.s.l.; southward, the bar is limited by another through deep up to -8 m a.s.l. Close to the northern harbour breakwater, a more extended surf zone occurs, about 160 m long between -4.5 and -5.5 m a.s.l. Interesting features can be observed

southward from the pier, where a northwest-southeast bar, high up to 5 m (from -9 to -4 m a.s.l.) represents the prolongation of the same pier. The bar is separated from the coastline by a deep up to -9 m a.s.l. In the southern sector, we observed a more extended surf zone, about 150 m long between -4.5 and -5.5 m a.s.l.

The marine weather in the physiographic unit, including the Cetraro harbour, is directly exposed to wind action from southwest, West and North-West across the Tyrrhenian Sea. The dominant wind blows from the west-southwest direction with a geographical fetch greater than 1650 km and speeds up to 7 Beaufort ( $\geq 28$  knots). They are frequent in winter, producing the highest sea waves up to more than 5 m. The winds from the northwest are prevailing winds with a maximum fetch of 650 km and speeds less than 6 Beaufort ( $< 22$  knots). They blow mainly in the spring and summer seasons. The longshore flow of sediments, moving from north to south, is interrupted by the harbour dock.

The data acquired by the Cetraro Buoy show a significant wave height of 0.65 m with an average period of 5.15 s. Inshore, the wave climate calculated is characterised by an average energy flux of 6.1 kN/s and a significant wave height of 1.01 m with an average period of 5.7 s.

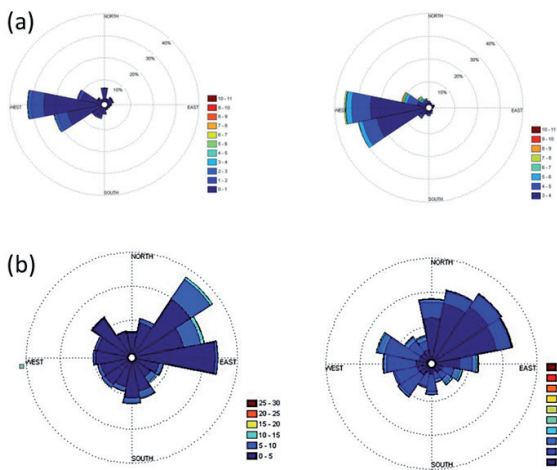


Fig. 5 - Frequency distribution of wave height and mean direction. (b) Wind behavior at the Paola (left rose diagram) and Belvedere (right rose diagram) stations

The application of EPM model to estimate the potential sediments yield of the Aron River catchment was based on bibliographic data (geological and landslide maps, thermal-pluviometric data, DEM). However, some coefficients included in the EPM calculation were calibrated by in situ measurements. In detail, the  $Y$  coefficient of the EPM, depending by the rock and soil resistance, was assigned to the different lithotypes cropping out in the catchment area based on 9 geomechanical surveys. During the geomechanical surveys, realized on the rock masses of the different

lithotypes, the dips, dip directions and strikes, spacing, length, aperture, and joint roughness coefficient (JRC) of discontinuities were measured. Furthermore, the Schmidt hammer rebound value was measured, and the Geological Strength Index (GSI - HOEK & BROWN, 1997) was estimated. Qualitative observations about the weathering of different lithotypes were realized.

At this step, the calculation with the EPM was carried out by keeping all the coefficients fixed ( $Y, y, Im$ , etc.) and making only the  $X$  coefficient (depending on land use) vary. The first calculation was performed using the Corine Land Cover 2012 land use map for the assignment of  $X$  coefficient. A second calculation was realized with the values assigned to the new land use classes coming from the processing of multispectral data (Land-Use 2023). In both cases, the values of the coefficients were assigned as proposed by ZEMLJIC, 1971. The different types of land use were redefined by means of multispectral analysis techniques. In detail, we used a classifier such as Random Forest, which performs well for multi-source classification of geographic and remote sensing data (GISLASON *et alii*, 2006).

In the first calculation, the value of  $W$  (average annual specific production of sediments) is equal to 29900  $m^3/year$ , while in the second one, it is equal to approximately 23600  $m^3/year$ . The difference between the values can be related to the degree of detail in defining the polygons with which the two land use classifications were created. In the case of the Corine Land Cover 2012, approximately 90 polygons differentiated into 10 main classes were identified in the Aron River. In the land use processed from multispectral data, more than 400 polygons divided into 18 classes were identified. It follows that the error in assigning coefficients to individual classes decreases greatly because each coefficient is applied to extremely homogeneous areas in terms of land use. Furthermore, comparing the two land use maps, it is possible to evaluate a decrease in the cultivated areas between 2012 and 2023, and consequently, an increase in the wooded areas.

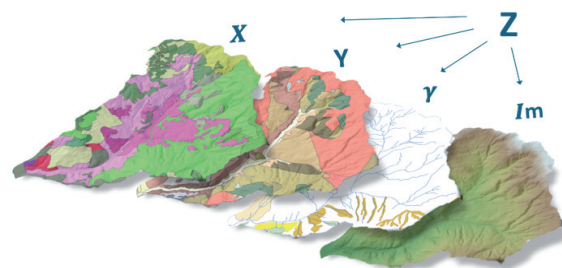


Fig. 6 - EPM coefficients relate to soil-use ( $X$ ), lithology ( $Y$ ) and geomorphology ( $\gamma$  and  $Im$ )

The results of this preliminary study on the impact of land use in estimating sediment production highlight some extremely interesting aspects on which further research can be carried out:

BY-PASS SEDIMENT SYSTEM TO REPLACE LONGSHORE CURRENT IN HARBOUR AREAS

Station	Date	Depth	Temperature	Conductivity	Salinity	Oxygen (O <sub>2</sub> (Sat))	Oxygen (O <sub>2</sub> (Data))	pH	Eh	Turbidity
S1	May-2023	0.75	21.7468	56.1068	38.1431	97.9	6.61	8.0095	407.2	66.156
		1	21.7471	56.1061	38.1423	97.9	6.62	8.0783	408.9	63.839
		1.75	21.75	56.1123	38.1436	97.9	6.62	8.0773	409.5	65.772
		2.25	21.7535	56.0807	38.1408	97.9	6.62	8.0765	410.6	65.998
		2.75	21.7515	56.0872	38.1409	97.9	6.62	8.0755	411.7	65.966
		3.25	21.7522	56.0911	38.1433	98	6.62	8.0733	412.8	63.979
		3.75	21.7396	56.0999	38.1445	98	6.62	8.0723	413.8	61.334
		4.25	21.7403	56.1028	38.1431	98	6.62	8.0715	414.5	60.325
		4.75	21.7409	56.1014	38.1433	98	6.62	8.0714	414.9	60.114
		5.25	21.7417	56.1001	38.1416	98	6.62	8.0633	415.4	58.348
		5.75	21.7426	56.1022	38.1423	98	6.62	8.0665	416.3	57.315
		6.25	21.7434	56.1004	38.1441	98	6.63	8.0683	417.5	55.826
		6.75	21.7434	56.1006	38.1418	98	6.63	8.0675	417.7	55.694
		7.25	21.7407	56.1	38.142	98	6.62	8.0655	418.4	54.636
		7.75	21.74	56.1018	38.144	98.1	6.63	8.0665	419.3	52.718
		8.25	21.7396	56.0995	38.1423	98.1	6.63	8.0641	419.9	51.309
		8.75	21.7396	56.0996	38.1437	98	6.63	8.0653	420.3	49.395
		9.25	21.7377	56.0964	38.1413	98	6.63	8.0611	420.7	48.615
		9.75	21.7363	56.0954	38.1403	98.1	6.63	8.0605	421.1	47.346
		10.25	21.7374	56.0907	38.1394	98.1	6.63	8.0605	421.5	47.189
		10.75	21.7322	56.0829	38.1385	98.2	6.64	8.0573	421.1	46.941
		11.25	21.7335	56.0796	38.1319	98.3	6.64	8.0573	421	46.182
		11.75	21.7328	56.0854	38.1174	98.3	6.64	8.0533	424	45.906
		12.25	21.7312	55.9591	37.9265	100.5	7.57	7.704	208.6	25.924
		12.75	21.7155	55.8513	37.9386	100.2	6.96	7.69	203.4	21.387
13.25	21.7184	55.8569	37.9376	99.2	6.71	7.826	197.4	19.925		
13.75	21.7217	55.8688	37.9417	99	6.7	7.831	196.3	19.136		
14.25	21.7214	55.8735	37.9428	99	6.7	7.847	196.1	19.139		
14.75	21.7219	55.8839	37.9469	99.2	6.71	7.845	194.5	19.832		
15.25	21.7189	55.8897	37.9499	99.2	6.72	7.854	194.9	19.744		
15.75	21.7215	55.9234	37.9499	99.4	6.72	7.865	194.8	19.782		
16.25	21.7265	55.9263	37.9684	99.4	6.72	7.865	194.8	19.782		
16.75	21.7223	55.90109	37.9676	99.5	6.73	7.881	193.1	19.354		
17.25	21.7211	55.9123	37.9688	99.4	6.72	7.912	191.9	19.603		
17.75	21.7167	55.9109	37.9699	99.4	6.72	7.914	191.9	19.603		
18.25	21.7188	55.9191	37.9723	99.4	6.72	7.922	191.9	19.587		
18.75	21.7206	55.92612	37.9742	99.4	6.72	7.911	191.6	19.184		
19.25	21.7294	55.92312	37.9792	99.5	6.73	7.928	191	18.711		
19.75	21.7321	55.9094	37.9757	99.5	6.73	7.954	190.6	18.885		
20.25	21.7312	55.9193	37.9762	99.6	6.73	7.943	190.3	17.943		
20.75	21.7201	55.92842	37.9873	99.2	6.73	7.958	190	17.831		
21.25	21.7192	55.9408	38.0014	99.5	6.73	7.954	190	17.711		
21.75	21.7162	55.9469	38.0045	99.4	6.73	7.973	189.9	17.142		
22.25	21.7178	55.9527	38.0015	99.5	6.73	7.959	189.9	17.101		
22.75	21.7122	55.95736	38.01416	99.6	6.73	7.962	189.9	17.137		
23.25	21.7201	55.96996	38.02117	99.6	6.73	7.964	189.9	17.148		
23.75	21.7122	55.9629	38.02686	99.5	6.73	7.965	189.8	16.987		
24.25	21.7261	55.9901	38.03138	99.5	6.73	7.964	189.8	16.142		
24.75	21.731	56.00862	38.04079	99.6	6.73	7.969	189.7	15.705		
25.25	21.7314	56.01881	38.0487	99.6	6.73	7.971	189.8	15.842		
25.75	21.7304	56.02894	38.0599	99.6	6.73	7.973	189.9	14.742		
26.25	21.7312	56.02877	38.05631	99.6	6.73	7.975	189.6	13.727		
26.75	21.7318	56.03946	38.06392	99.6	6.73	7.976	189.6	12.943		
27.25	21.7382	56.05358	38.07203	99.6	6.73	7.98	189.5	12.129		
27.75	21.7353	56.07331	38.08823	99.6	6.73	7.98	189.5	11.645		
28.25	21.7358	56.09916	38.09916	99.6	6.73	7.979	189.5	11.348		
28.75	21.7359	56.09914	38.10747	99.6	6.73	7.982	189.5	11.04		
29.25	21.7314	56.1128	38.1189	99.7	6.73	7.981	189.5	10.969		
29.75	21.7349	56.1287	38.1173	99.6	6.73	7.983	189.4	10.877		
30.25	21.7392	56.13939	38.1353	99.6	6.73	7.986	189.4	10.754		
30.75	21.741	56.1541	38.14502	99.7	6.73	7.986	189.4	10.668		
31.25	21.74412	56.1675	38.15245	99.7	6.73	7.986	189.4	9.444		
10.25	21.7441	56.18854	38.16994	99.7	6.73	7.987	189.4	9.214		
10.55	21.7465	56.21128	38.18207	99.7	6.72	8.009	189.5	9.094		

Tab. 1 - Results of the multiparametric logs in the point SM1 and SM2 during May 2023

- By applying the multispectral processing methods used in this study it is possible to carry out highly localized analyses of land use changes over time and quantitatively identify the impact of this variations on sediment supply. The time span analyzed in this study is very broad as the objective was to evaluate the impact of anthropic practices quantitatively and generally on the dynamics of sediment supply and transport at the basin scale;
- The quantitative estimation of the sedimentary deficit and the identification of the areas where this is most marked allows to objectively plan specific interventions;
- The possibility of carrying out these procedures on a large scale allows, after having calibrated the model, to expand the calculation to larger areas and evaluate at the scale of a physiographic unit the variation in sedimentary supply and the consequent relationship with the variations in erosion rates in coastal areas.

In the investigated coastal sector, the littoral transport, moving from north to south, is characterised by an average value of about 80,000 m<sup>3</sup>/yr. Part of this amount is trapped at the harbour mouth (about 10,000 m<sup>3</sup>/yr) and in the longitudinal bar southward from the pier (about 50,000 m<sup>3</sup>). These trapped sediments, combined with the low potential sediment supply by the Aron River

Station	Date	Depth	Temperature	Conductivity	Salinity	Oxygen (O <sub>2</sub> (Sat))	Oxygen (O <sub>2</sub> (Data))	pH	Eh	Turbidity
S1	October-2023	0.75	23.9319	56.1068	38.1696	97.9	6.61	8.105	409.9	1.351
		1	23.9319	56.1068	38.1696	97.9	6.61	8.105	409.9	1.351
		1.75	23.9322	56.1061	38.1688	97.9	6.62	8.103	411.6	1.054
		2.25	23.9391	56.1123	38.1701	97.9	6.62	8.102	412.3	0.967
		2.75	23.9364	56.0805	38.1673	97.9	6.62	8.101	413.3	0.293
		3.25	23.9166	56.0872	38.1674	97.9	6.62	8.1	414.4	0.761
		3.75	23.9173	56.0911	38.1698	98	6.62	8.098	415.5	0.774
		4.25	23.9277	56.0999	38.1711	98	6.62	8.097	416.5	0.709
		4.75	23.9254	56.1028	38.1716	98	6.62	8.096	417.3	0.732
		5.25	23.9209	56.1013	38.171	98	6.62	8.095	417.6	0.709
		5.75	23.9268	56.1001	38.1691	98	6.62	8.095	418.1	0.743
		6.25	23.9277	56.1022	38.1688	98	6.62	8.094	419	0.71
		6.75	23.9287	56.1054	38.1706	98	6.63	8.093	419.9	0.821
		7.25	23.9275	56.1009	38.1678	98	6.63	8.092	420.4	0.889
		7.75	23.9238	56.1	38.1685	98	6.62	8.092	421.1	0.821
		8.25	23.9251	56.1018	38.1705	98.1	6.63	8.091	422	0.773
		8.75	23.9249	56.0993	38.1686	98.1	6.63	8.089	422.6	1.485
		9.25	23.9247	56.0996	38.1662	98	6.63	8.088	423	0.79
		9.75	23.9238	56.0964	38.168	98	6.63	8.086	423.8	0.81
		10.25	23.9214	56.0994	38.1688	98.1	6.63	8.085	423.8	1.341
		10.75	23.9191	56.0897	38.1659	98.1	6.63	8.082	424.2	2.379
		11.25	23.9185	56.0829	38.163	98.2	6.64	8.082	424.8	1.196
		11.75	23.9186	56.0796	38.1594	98.3	6.64	8.082	425.7	1.577
		12.25	23.9209	56.0824	38.1499	98.2	6.64	8.083	426.7	2.101
		12.75	23.9803	56.1074	38.1286	98.7	6.66	8.075	190.1	0.331
13.25	23.9818	56.1148	38.133	98.7	6.66	8.077	190.1	0.357		
13.75	23.9861	56.116	38.1301	98.7	6.66	8.085	190.2	0.288		
14.25	23.9919	56.1333	38.1382	98.7	6.66	8.085	190.1	0.378		
14.75	23.9961	56.1425	38.1417	98.8	6.67	8.085	190	0.357		
15.25	23.9817	56.1506	38.1438	98.8	6.67	8.081	190	0.425		
15.75	23.9816	56.1513	38.1435	98.8	6.67	8.081	190	0.299		
16.25	23.9803	56.1513	38.1437	98.8	6.67	8.089	189.9	0.312		
16.75	23.9936	56.1421	38.1431	98.8	6.67	8.057	189.9	0.394		
17.25	23.9876	56.1323	38.1408	98.8	6.67	8.055	189.9	0.35		
17.75	23.9875	56.1356	38.1432	98.8	6.67	8.056	189.8	0.274		
18.25	23.9894	56.1381	38.1434	98.8	6.67	8.056	189.8	0.349		
18.75	23.9843	56.1353	38.1434	98.8	6.67	8.055	189.8	0.388		
19.25	23.9839	56.1329	38.1404	98.8	6.67	8.051	189.8	0.335		
19.75	23.9911	56.1434	38.1437	98.9	6.68	8.051	189.8	0.374		
20.25	23.9934	56.146	38.1436	99	6.68	8.051	189.8	0.338		
20.75	23.9958	56.1484	38.1433	99	6.68	8.047	189.8	0.268		
2										

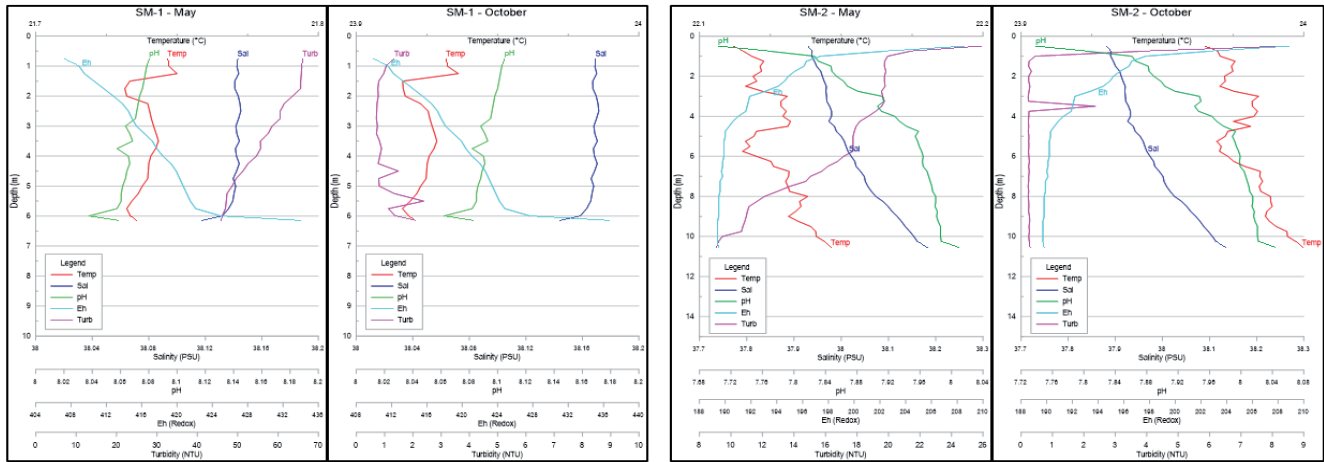


Fig. 7 - May and October multiparametric logs in the point SM1 and SM2

sample	Sediment type	distribution	gravel					sand					D10 (µm)	D50 (µm)	D90 (µm)
			very coarse	coarse	medium	fine	very fine	very coarse	coarse	medium	fine	very fine			
S1	Sandy Medium Gravel	Bimodal, Poorly Sorted	0%	0%	28,4%	13,9%	18,2%	24%	12,1%	3%	0,4%	0%	771,2	2876,5	9948,5
S2	Sandy Fine Gravel	Bimodal, Poorly Sorted	0%	0%	7%	23%	21,9%	26,6%	14,6%	5,8%	0,8%	0,2%	601,4	2127	7021,1
S3	Gravelly Very Coarse Sand	Moderately Sorted	0%	0%	1,5%	3,7%	14,6%	54,5%	18%	3,7%	3%	1%	591,1	1323,4	2798,8
S4	Sandy Very Fine Gravel	Trimodal, Poorly Sorted	0%	0%	17,6%	18,5%	20,1%	25,3%	13,3%	4,4%	0,6%	0,1%	684,1	2448,4	9252,8

Tab. 3 - Results of the grain size analysis

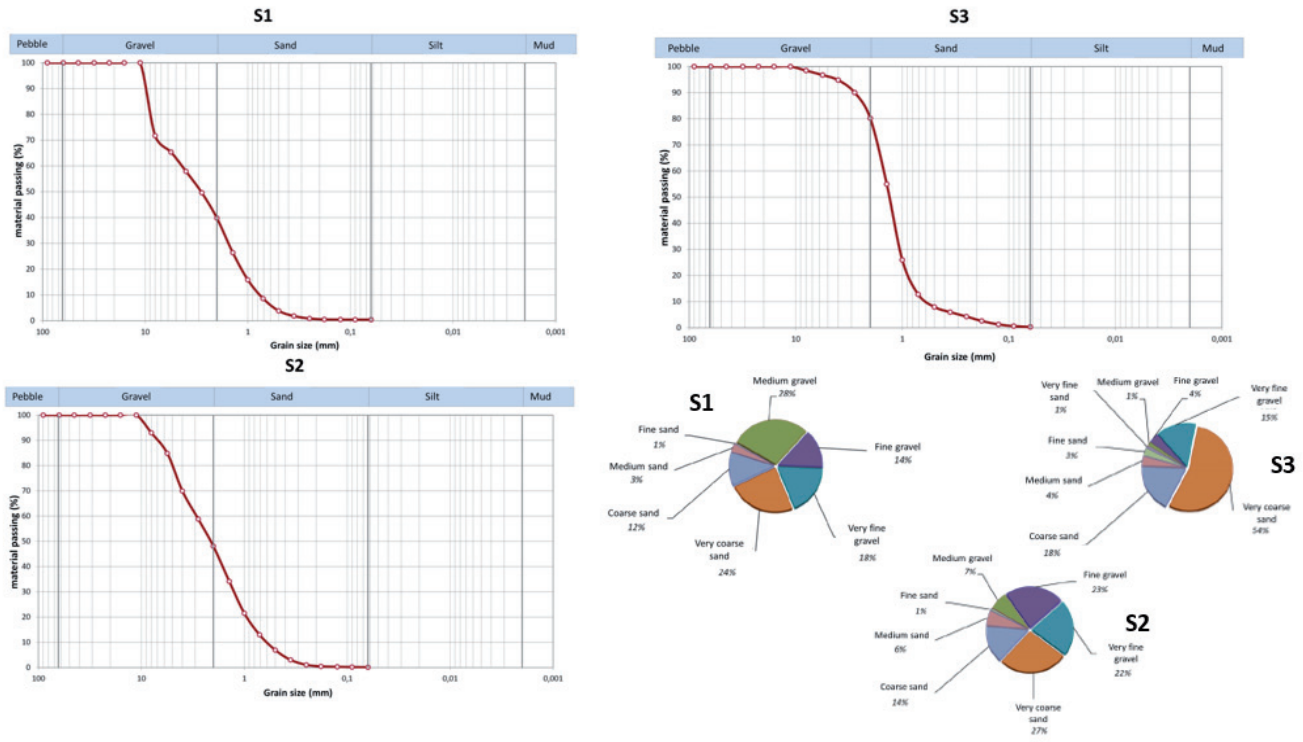


Fig. 8 - Grain size distribution curves



During the survey of October 2023 we obtained the following results for the point SM1 (Table 2): (i) temperature of 23.9 °C; (ii) electrical conductivity of about 56 mS/cm; (iii) salinity around 38.5 PSU; (iv) dissolved oxygen concentration of about 6.6 ppm at top and bottom respectively; (v) pH of 8.1; (vi) Eh varying from 409 mV to 436 mV from top to bottom; (vii) turbidity less than 1 NTU. For the point SM2 the results are: (i) temperature of 23.9 °C; (ii) electrical conductivity of about 56 mS/cm; (iii) salinity around 38.1 PSU; (iv) dissolved oxygen concentration of about 6.6 ppm (v) pH of 8; (vi) Eh varying from 189.8 mV to 190.1 mV from top to bottom; (vii) turbidity less than 1 NTU.

The seawater column monitoring is strongly influenced by wave climate and environmental conditions of the monitoring points. An example is represented by the fluvial plume. Indeed, the suspended sediment discharged in front of the fluvial mouth can influence many seawater column parameters, particularly turbidity. The effects of fluvial plumes on the seawater turbidity are widely analysed in the recent literature, especially thanks to the approaches based on remote sensing. For example, TAVORA *et alii* (2023) developed an algorithm, calibrated with in situ measurements, to calculate the turbidity from satellite images; the authors tested the algorithm in the Patos Lagoon (Brazil), obtaining seawater turbidity values, related to fluvial plumes, up to 50 NTU; COVELLI *et alii* (2007) investigated the effects of fluvial plumes on the physico-chemical parameters of seawater in the northern Adriatic Sea (Italy), recording turbidity values, after a river flood, of 50-60 NTU.

In the study area, the turbidity value of up to 60 NTU recorded during May 2023 can be related to the fluvial plume produced by Aron River flood.

The coastal sediment samples were collected at four different points (Fig. 4). In particular, S1 and S2 were sampled in correspondence with the emerged sediment accumulation close to the harbour mouth, while S3 in the submerged area. S4 was sampled in the Lampetia area, northward and up-drift than the Cetraro harbour. The grain-size analysis showed the following features: (i) S1 is a poorly sorted sandy medium gravel; (ii) S2 is poorly sorted sandy fine gravel; (iii) S3 is moderately sorted very fine gravelly/very coarse sand; (iv) S4 is poorly sorted sandy very fine gravel.

The results of the grain-size analysis confirm the prevalence in the study area of coarse-grained sediments reported in literature (CNR, 1997). Furthermore, the new data show that the sediments of emerged area are coarse grained and poorly sorted (bimodal S1 and S2, even trimodal S4), index of poor sediments reworking by waves. Instead, the sediments of submerged area (S3) are reworked by waves as showed by moderated sorting.

On the same 4 sediment samples, ecotoxicological analysis were realized. In detail, the toxicity of the solid fraction of samples (excluding the particles >5 mm of biological, geological, and anthropic origin) was tested by the use of *V.fischieri* marine bacteria, according to APAT-ICRAM (2007).

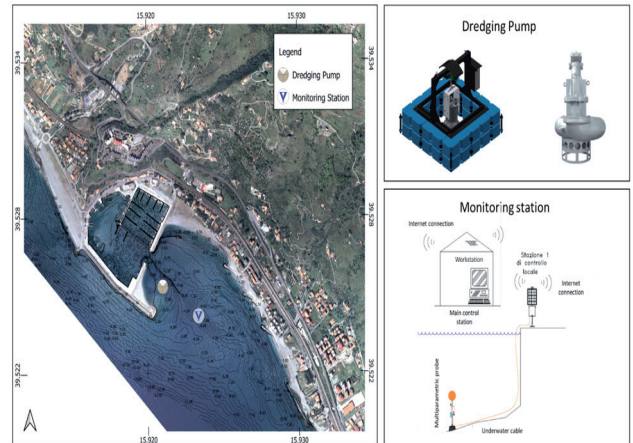


Fig. 9 - Scheme of the bypass and monitoring system

The toxicity of all 4 samples resulted absent/negligible.

Collected data combined with the multidisciplinary approach allowed to design an innovative and sustainable technological solution permitting to contrast the mouth silting and coastline erosion in the down-drift sector.

The beach nourishment in the down drift area with respect to the harbour structure can be realised by means of by-pass systems, which is a low-impact operations recovering the sediment loss due to the coastal dynamic imbalance. The nourishment via a by-pass system normally lasts few days, whereas traditional nourishments using tracks to move the sediment can last longer, implying multiple environmental stresses. A by-pass system implemented with a fixed dredge pump involves the use of a specific electric pump for dredging, properly dimensioned on the distance to be reached. The pump's sizing must ensure the correct velocity of the mixture inside the pipeline, in order to avoid sand settling along the pipeline. An electrical panel controls the pump, which can be manually or timed controlled, by setting the frequency and duration of the pump's operation. A sand pipeline, adequately sized to transport the mixture to the identified accumulation zone is connected to the pump. Being a fixed installation, the pipeline can be buried to avoid interference with maritime traffic.

## CONCLUDING REMARKS

The sedimentary balance of the watershed-coast-continuum scale is very important in managing the fluvial and coastal areas. In the Tech4You project (T4Y S2G2PP2), we focused on the silting process affecting the Cetraro harbour mouth, a common problem of a lot of harbours realised in high wave energy areas. In fact, the realization of marine infrastructures usually modifies the wave approach to the coast as well as the longshore current and the sediment drift regime. Usually, the dredging process, involving the removal of sediment in its natural deposited condition by using either mechanical or hydraulic equipment, is carried out. However,

dredging implies relevant drawbacks, such as strong impacts on marine flora and fauna mobilization and diffusion of contaminants and pollutants. The aim was to define an innovative and sustainable technological solution to reduce costs and environmental impacts. We characterised the physiographic unit in terms of wave climate and currents and tried to refine the estimation of sedimentary balance. Also the morphology of emerged and submerged areas characterising the sediment's grain size was analysed, and the main physical-chemical parameters (turbidity, temperature, electrical conductivity, salinity, dissolved oxygen, pH, Eh) of the seawater column typically influenced by dredging operations, were monitored. Finally, a bypass system was designed, consisting of a submerged pump installed at the end of the pier on the up-drift side of the harbour. The bypass system allows the movements of the sediments from the trap, created

by the harbour structure, to the down-drift coastline. Hence, this system aims to replace the original longshore current in the section of the coast where the littoral current has been interrupted since the harbour building. Simultaneously, the bypass system allows the minimization of the environmental drawbacks produced by the classical dredging operations.

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